

Nicaraguan volcanoes record paleoceanographic changes accompanying closure of the Panama gateway

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¹GSA Data Repository item 2002#####, Age data, geochemistry of sediments and volcanic rocks, and bulk sediment flux calculations, is available from Documents Secretary, GSA, P.O. Box 9140, Boulder, CO 80301-9140, editing@geosociety.org, or at www.geosociety.org/pubs/ft2002.htm.

ABSTRACT

A major oceanographic event preserved in the Cocos plate sedimentary column has survived subduction and been recorded in the changing composition of Nicaraguan magmas. A uranium increase in these magmas since the latest Miocene (after 7 Ma) resulted from the “carbonate crash” at 10 Ma and the ensuing high organic carbon burial in the sediments. The response of the arc to this paleoceanographic event requires near steady-state sediment recycling at this margin since 20 Ma. This relative stability in sediment subduction invites one of the first attempts to balance sedimentary input and arc

output across a subduction zone. Calculations based on Th indicate that as much as 75% of the sedimentary column has been subducted beneath the arc. The Nicaraguan margin is one of the few places to observe such strong linkages between the oceans and the solid earth.

INTRODUCTION

The growth rate of the continental crust and chemical evolution of the mantle depend in large part on the balance between subducted input and volcanic output at subduction zones. Sediment subduction represents a loss of mass from the continents and a gain of exotic chemical components in the mantle. The recycling of sedimentary material back to the continents via arc magmatism, however, modulates these effects. Despite the importance to the evolution of the Earth, few studies have attempted to balance mass or chemical components across subduction zones (Patino et al., 2000; Bach et al., 1998). Many difficulties arise in attempting such a mass balance. For example, marine sections near trenches have been well sampled by ocean drilling (e.g., Rea and Ruff, 1996, Plank and Langmuir, 1998), but several million years of subduction separates these materials from those currently contributing to the magma-genesis zone, typically >100 km beneath the arc. This delay of a few million years is inherent to the recycling problem. Can we make sensible predictions about the material flux to depth, and/or can we assume that some aspects of the system are in steady state?

Previous work on the temporal variation of Western Pacific arcs has found that although some may respond to tectonic events such as backarc spreading (e.g., Tonga; Clift and Vroon, 1996), others have maintained a relatively constant geochemical composition for millions of years (e.g., Izu; Bryant et al., 1999). The sediment-recycling signature in these arcs, however, is weak and thus not ideal for exploring the temporal variability of sediment flux and delivery. On the other hand, sediment tracers such as the ^{10}Be abundance and the Ba/La ratio reach their global maximum in northwestern Nicaragua (Carr et al., 1990; Tera et al., 1986). Nicaragua also has the advantage that the arc has migrated trenchward with time, and so unlike other arcs where the volcanic history is buried by each successive eruption, the past >20 m.y. of arc volcanism is exposed in surface outcrops in Nicaragua (Ehrenborg, 1996).

TEMPORAL VARIATIONS IN COCOS PLATE SEDIMENTS

The other advantage to studying the temporal response to sediment subduction in Nicaragua is the simple sedimentary stratigraphy on the incoming Cocos plate (Fig. 1). Results from Deep Sea Drilling Project (DSDP) Leg 67 and Ocean Drilling Program (ODP) Legs 138 and 170 have provided a detailed history of sedimentation in the Guatemala Basin (Aubouin et al., 1982; Mayer et al., 1992; Kimura et al., 1997). The sedimentary column consists largely of two units: pelagic carbonate overlain by diatomaceous hemipelagic clayey ooze (Fig. 2). The boundary between these two units is roughly synchronous across the basin at ca. 10 Ma and represents the “carbonate crash” (Lyle et al., 1995). The crash is thought to be caused by an ~800 m rise in the carbonate compensation depth, apparently reflecting the cessation of flow of less corrosive Caribbean intermediate and deep water into the Pacific following the shoaling of the Panama sill (Lyle et al., 1995; Farrell et al., 1995). Postcrash sediments were enriched in organic carbon as the seafloor approached the Costa Rica dome, a region of very high open-ocean surface productivity (Hofmann et al., 1981).

Because geochemical tracers are the means by which we detect sediment recycling, we need to relate the stratigraphic changes to changes in the geochemical fluxes into the trench. Toward this end, we analyzed¹ 16 samples from ODP Sites 844 and 845 (Fig. 2), which, when combined with data from DSDP Site 495 (Patino et al., 2000), provide a chemical stratigraphy of Guatemala Basin sediments. Although most elements have very different concentrations in the two lithologic units, we focus here on Ba and U, which are notably enriched in the modern Nicaraguan volcanic arc, where they correlate with unambiguous sediment tracers such as ¹⁰Be.

U abundance varies with lithologic type, being low in the lower pelagic carbonate unit (<500 ppb), but high in the upper hemipelagic ooze unit (up to 10 ppm). As is commonly observed, U follows organic C owing to authigenic precipitation of reduced U(IV) oxides (Chase et al., 2001; Klinkhammer and Palmer, 1991). At face value, the steep increase in both organic C and U in the top 25 m of the columns might be expected as any site approaches land, receives terrestrial carbon, and then suffers diagenetic release. There are several reasons, however, why this generic model does not appear to

be the dominant control on the U and C distributions. For one, these sites are at > 3000 m water depth and far from coastal upwelling regimes and terrestrial carbon sources. This is consistent with C/N ratios in Site 844 and 845 sediments, which indicate predominantly marine sources (Meyers, 1997). The lack of sulfate reduction or ammonium production in pore fluids at ODP Sites 844 and 845 (Mayer et al., 1992) suggests that diagenesis may not be the major controlling process. Instead, the increase of U and organic C during the past few million years appears to coincide with passage of the sites beneath the increased productivity caused by upwelling in the Costa Rica dome (Fig. 1a). Although there is little direct information on the history of the Costa Rica dome, it likely strengthened at ca. 5–4 Ma, when the shoaling of the Panama sill caused major changes in basin salinity (Haug et al., 2001), thermocline depth (Cannariato and Ravelo, 1997), surface currents (Nisancioglu et al., 2002), and position of the intertropical convergence zone (Hovan, 1995); the latter may have the largest effect on the formation and position of the dome (Hofmann et al., 1981).

Although high U at the top of the section seems to reflect high productivity in the Costa Rica dome, the low U abundances within the lower carbonate-rich unit are most likely due to a lack of preservation of organic matter. Carbonate-rich sediments typically have lower organic C abundances than clay-rich sediments. The reasons are debated, but the lower surface area of carbonate sediments, owing in part to the abundance of sand-sized foraminifera, may lead to poorly preserved organic C (Mayer, 1994; Pedersen, 1995) and by association, U. For example, ODP Sites 845 and 850 both underlie regions with similar modern surface productivity (Lyle, 1992), and yet the carbonate-rich surface sediments at Site 850 contain much lower organic C (<0.4 wt%) than the higher-surface-area clay-rich surface sediments at Site 845 (>2 wt% organic C). Thus, the U profile in the Cocos sediments reflects both biologic productivity and preservation. The closure of the Panama gateway led to the carbonate crash at ca. 10 Ma and strengthened the Costa Rica dome at ca. 5–4 Ma, both of which enhanced precipitation, preservation, and subduction of U in sediments at the Central American Trench.

In contrast to U, Ba accumulation rates were high prior to the carbonate crash and have decreased somewhat since. The entire column preserves a high proportion (>95%) of authigenic Ba (Schroeder et al., 1997), and the lithologic changes in the section appear

to have had little effect on Ba preservation. High Ba concentration (>2000 ppm) is a long-term (>16 m.y.) feature of the subducting sedimentary column, whereas high U is not. Because mass-accumulation rates are directly proportional to subduction-input rates, high Ba accumulation translates into high Ba flux to the Central American Trench throughout the past 20 m.y., whereas the U flux has increased dramatically during the past several million years.

By using the foregoing inferences, we calculate sediment-subduction fluxes at 12 Ma, just prior to the carbonate crash, and at 2.5 Ma, which is the age of material currently beneath the arc (Table 1). These calculations show that the sedimentary flux of U to the trench has increased by a factor of 4, whereas the flux of Ba has varied only by ~10%.

SEDIMENT RECYCLING AT THE NICARAGUA ARC

In order to gauge the response of the Nicaragua arc to these temporal changes in the sediment input, we conducted a field campaign in 1996 to sample the Miocene Coyol arc (Ehrenborg, 1996; Fig. 1). Ar-Ar dates (see footnote 1) range from 24 to 7 Ma, and the arc has migrated trenchward since ca. 13 Ma (Fig. 1). The unsampled 7–0 Ma volcanic section is presumably buried beneath volcanoclastic deposits in the Nicaragua depression. We collected ~60 samples (Balzer, 1999) and analyzed ~40 of the most mafic and least altered for major and trace elements and Sr isotopes (see footnote 1). In order to expand our coverage, we analyzed another ~20 samples collected previously (Nystrom et al., 1988). The close similarities between the Coyol and modern arcs in major element and rare earth element compositions (Balzer, 1999) suggest that the Coyol volcanic rocks most likely represent the main Miocene arc, and not backarc volcanism.

Remarkably, the Coyol arc preserves the same along-strike trend in the Ba/La ratio as the modern arc, both in magnitude and spatial gradient (Fig. 3). This coincidence suggests that both the incoming flux of Ba and the delivery process have been fairly constant along this margin for 20 Ma. On the other hand, the Coyol arc is uniformly depleted in U relative to the modern arc (Fig. 3), a finding consistent with predicted stratigraphic changes in Cocos plate sediments. Because all the Coyol samples are older than 7 Ma and because the subduction time is on the order of 2.5 Ma, the sedimentary columns contributing to the Coyol arc all preceded the carbonate crash. In addition to U

and Ba, other chemical systematics are consistent with the predicted sedimentary fluxes. The Coyol and modern arcs are similar in their range of Ce/Pb ratios and Ce anomalies, as are the predicted sediment columns, whereas the Coyol arc has slightly lower $^{87}\text{Sr}/^{86}\text{Sr}$ ratios, as predicted from the global seawater Sr isotope curve. Thus, the temporal stability (in Ba, Ce/Pb, REEs) and changes (in U and Sr isotopes) in the Nicaraguan arcs appear to originate in the oceanographic processes that control Cocos plate sedimentation.

Given the regularity in the subduction process at Nicaragua, it is possible to balance the input and output fluxes for some element tracers. A full mass balance would require an estimate of the composition of the subducting basaltic basement as well, as this layer also contributes elements to the arc (e.g., Elliott, et al., 1997). In fact, basaltic U buffers arc concentrations, which is the reason that the modern and Miocene arcs vary by at most a factor of 2 (Fig. 3), whereas the sediment flux varies by a factor of 4 (Table 1). Unfortunately, the Cocos basement has been poorly sampled in this region, so we focus here on the elements Th and Ba, whose budgets in this arc are dominated by sediment input (Plank and Langmuir, 1993). We calculate (Table 1) that 76% of the sedimentary Th and 28% of the sedimentary Ba are recycled to the arc. These results are for the average Nicaraguan arc and do not take into account the large gradient in Ba/La ratio along the arc, which reflects important along-strike variations in delivery, the source of which is debated (Carr et al., 1990; Leeman et al., 1994; Herrstrom et al., 1995; von Huene et al., 2000). We obtain virtually the same mass balance for the 12 Ma sedimentary column and average Miocene arc (80% of the Th and 28% of the Ba), assuming a constant material flux for the arc. These calculations demonstrate near-steady-state recycling efficiency of this margin since 20 Ma. The preferential recycling of Th over Ba is not predicted from partition coefficients for red clay (Johnson and Plank, 1999), but may have to do with the different behavior of the primary Ba host (sulfate) and Th host (silicate) in the subduction zone, or the shallow loss of Ba to a fluid during sulfate reduction (Shipboard Scientific Party, 1997). Regardless of the exact mechanisms, the magnitude of Th recycling requires subduction of at least 75% of the sedimentary column beneath the arc. This result is consistent with efficient ^{10}Be recycling (Tera et al., 1986), as well as the apparent lack of sedimentary accretion at the Nicaragua margin (Ranero and von Huene, 2000).

The fidelity of the recycling process at the Nicaragua "subduction factory" illustrates how tectonic events reverberate through the Earth. Tectonic uplift of the Isthmus of Panama restricted and ultimately ended communication between the tropical Atlantic and Pacific Oceans. This episode caused changes in seawater chemistry and ensuing sedimentation, which then affected element fluxes to the Central American Trench and ultimately the composition of magmas erupted in the volcanic arc. Although Patino et al. (2000) have shown how the two sedimentary units on the Cocos plate can be resolved in eruptive products from a single volcano, we show how secular variations in the two units have been recorded in the arc during the past 20 m.y. It remains to be seen whether we can extend this treatment to volatile components and assess the impact of the carbonate crash and subduction cycling on the long-term global carbonate-CO₂ cycle.

ACKNOWLEDGMENTS

We thank J-O. Nyström for generously providing Coyol arc samples and Wilfried Strauch and Nelson Buitrago at INETER (Instituto Nicaraguense de Estudios Territoriales) for help in Nicaragua. Rick Murray, Maureen Raymo, Andy Kurtz, Christina Ravelo, Simon Turner and Mitch Lyle are thanked for discussions and reviews and Dan Gravatt for technical assistance. We thank Bruce Idleman at Lehigh University for performing Ar-Ar dates. Eli Silver, Kirk McIntosh, and Julie Morris were essential in the initiation of this project. We acknowledge support from the following U.S. National Science Foundation grants: OCE-9521717 (to Plank) and EAR-9628251, EAR-9406624, and OCE-9521716 (to Carr).

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FIGURE CAPTIONS

Figure 1. A: Central American volcanic arc (triangles) and Guatemala Basin (drill sites shown). Arrows indicate ocean circulation (North Equatorial Countercurrent). Dotted circle is high-productivity Costa Rica dome (drawn where thermocline shoals to 40 m in August–November, Fiedler et al., 1991). Shading is seafloor depth; contours labeled in meters. B: Volcanic arc in Nicaragua. Active volcanic front (solid triangles) and Miocene Coyol volcanic arc (open circles are samples dated by Ehrenborg [1996]; open squares are samples dated in this study). Age ranges given in Ma. Lago Nicaragua (LN) and Lago Managua (LM).

Figure 2. Litho- and chemostratigraphy at DSDP Site 495 and ODP Sites 844 and 845, Guatemala Basin. A and B: Calcium carbonate concentrations and organic carbon mass-accumulation rate downcore (Aubouin et al., 1982; Mayer et al., 1992). C and D: Ba and

U mass-accumulation rate downcore. Site 495 data from Patino et al. (2000); Sites 844 and 845 ICP-MS data in Data Repository (see footnote 1). Shaded bar is 11–9 Ma time period of carbonate crash in Guatemala Basin and separates carbonate-rich from hemipelagic unit. High organic C and U at top of sections is most likely due to passage of sites beneath productive Costa Rica dome. The higher U concentrations at Site 495 are consistent with its longer history beneath the dome.

Figure 3. Geochemical variations along Central American volcanic arc, illustrating similar Ba/La ratios but contrasting U/Th ratios between modern and Miocene volcanic rocks in Nicaragua sector. These relationships are consistent with changes calculated for sedimentary columns subducted at 2.5 Ma and 12 Ma in Central American Trench (Table 1), shown as arrows (some off-the-scale). Modern arc data from Carr and Rose (1987); Miocene data in Repository (see footnote 1). Samples plotted have <53% SiO₂ (modern) or >3% MgO (Miocene), and <2% loss on ignition (least altered), and exclude high TiO₂ basalts (as in Carr et al., 1990). Alteration studies demonstrate that severe visible alteration (more than that for samples plotted here) leads to <6.5% U loss and 17% Ba addition (Balzer, 1999). U/Th and Ba/La ratios are plotted instead of U and Ba concentrations in order to minimize effects of partial melting and crystal fractionation and to distinguish from mid-oceanic-ridge basalts (MORBs) and oceanic-island basalts (OIBs). Average arc concentrations (no screening or normalizing) also reflect similar Ba (within 7%) and markedly different U (50%) between modern and Miocene arcs. Relationship between arc and sediment ratios is not 1:1 because ratios are fractionated as sedimentary material leaves subducting plate and mixes with mantle beneath arc.

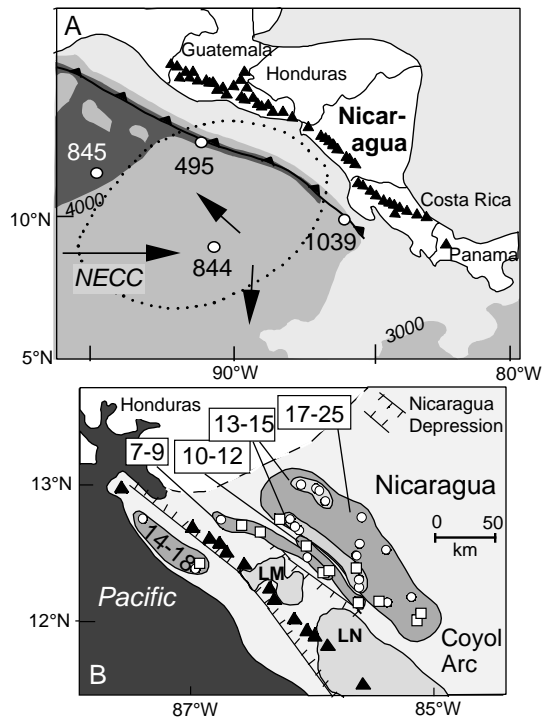


Figure 1. Plank, Balzer and Carr

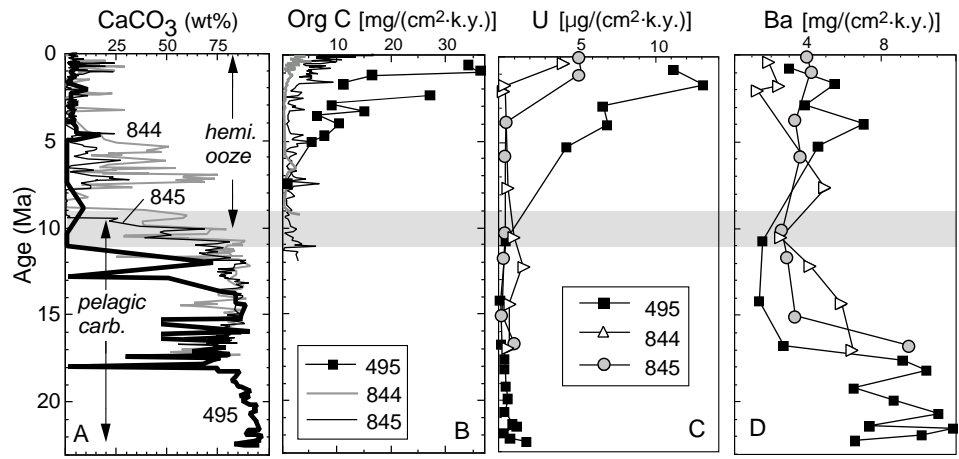


Figure 2. Plank, Balzer and Carr

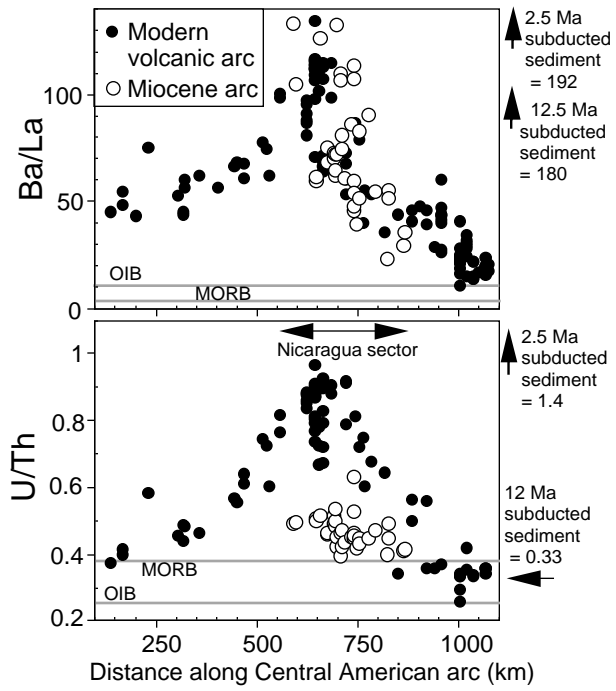


Figure 3. Plank, Balzer and Carr

TABLE 1. CONCENTRATIONS AND FLUXES

	U (ppm)	Th (ppm)	Ba (ppm)	La (ppm)	Material flux* ($\times 10^{10}$ kg/Ma)
Bulk sediment, 2.5 Ma	1.17	0.81	2705	14.0	2.8
Bulk sediment, 12 Ma	0.30	0.91	2422	13.5	2.9
Avg. Modern arc	0.69	0.89	663	6.9	3.3
Avg. Miocene arc	0.46	1.01	621	8.6	3.3
Fractionated MORB	0.14	0.37	28	7.8	

Sediment and Miocene arc concentrations are given in GSA Data Repository (see footnote 1 in text). Modern arc data from Carr and Rose (1987). Sediment column currently beneath Nicaragua was subducted at 2.5 Ma (on the basis of 68 mm/yr convergence rate, 68° slab dip, and 158 km depth to slab beneath arc). Arc flux calculated after subtracting mid-oceanic ridge basalt (2x N-MORB in Hofmann, 1988). Sediment flux at 2.5 Ma from 464 m column and 885.6 kg/m³ dry bulk density.

Sediment flux at 12 Ma from 454 m, 946.6 kg/m³. Volcanic output rates from Patino et al. (2000); 1.17×10^7 m³/Ma* and 2800 kg/m³ density.

* per meter of arc length